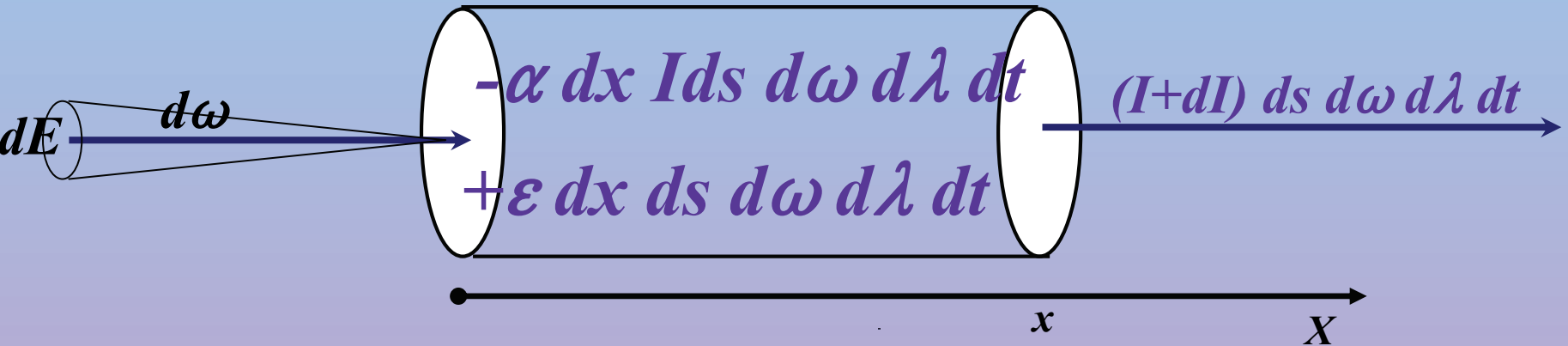


Basics of radiative transfer

Radiative transfer equation

$dE = I ds d\omega d\lambda dt$ – definition of intensity



$$\frac{dI}{dx} = -\alpha I + \epsilon$$

or

$$\frac{dI}{d\tau} = -I + S$$

α - volume extinction coefficient

ϵ - volume emission coefficient

$d\tau = \alpha dx$ – optical depth

$S = \epsilon / \alpha$ - source function

Basics of radiative transfer

1. General solution

$$I(x) = I_0 e^{-\int_0^x \alpha(x') dx'} + \int_0^x \mathcal{E}'(x') e^{-\int_{x'}^x \alpha(t) dt} dx'$$

2. Empty space ($\alpha=0$, $\varepsilon=0$)

$$I(x) = \text{const}$$

3. Medium without sources

($\alpha \neq 0$, $\varepsilon=0$) (Beer's law)

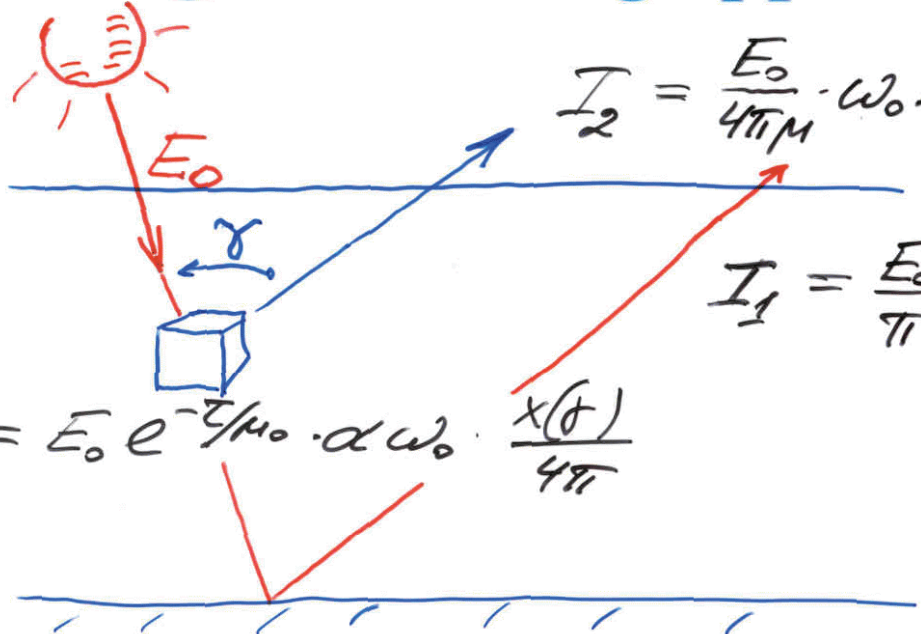
$$I(x) = I_0 e^{-\int_0^x \alpha(t) dt}$$

Scattered solar light in the atmosphere

1. Radiative transfer equation

$$\mu \frac{dI}{dx} = -\alpha I + \mathcal{E}$$

2. Single scattering approximation ($\tau \ll 1$)

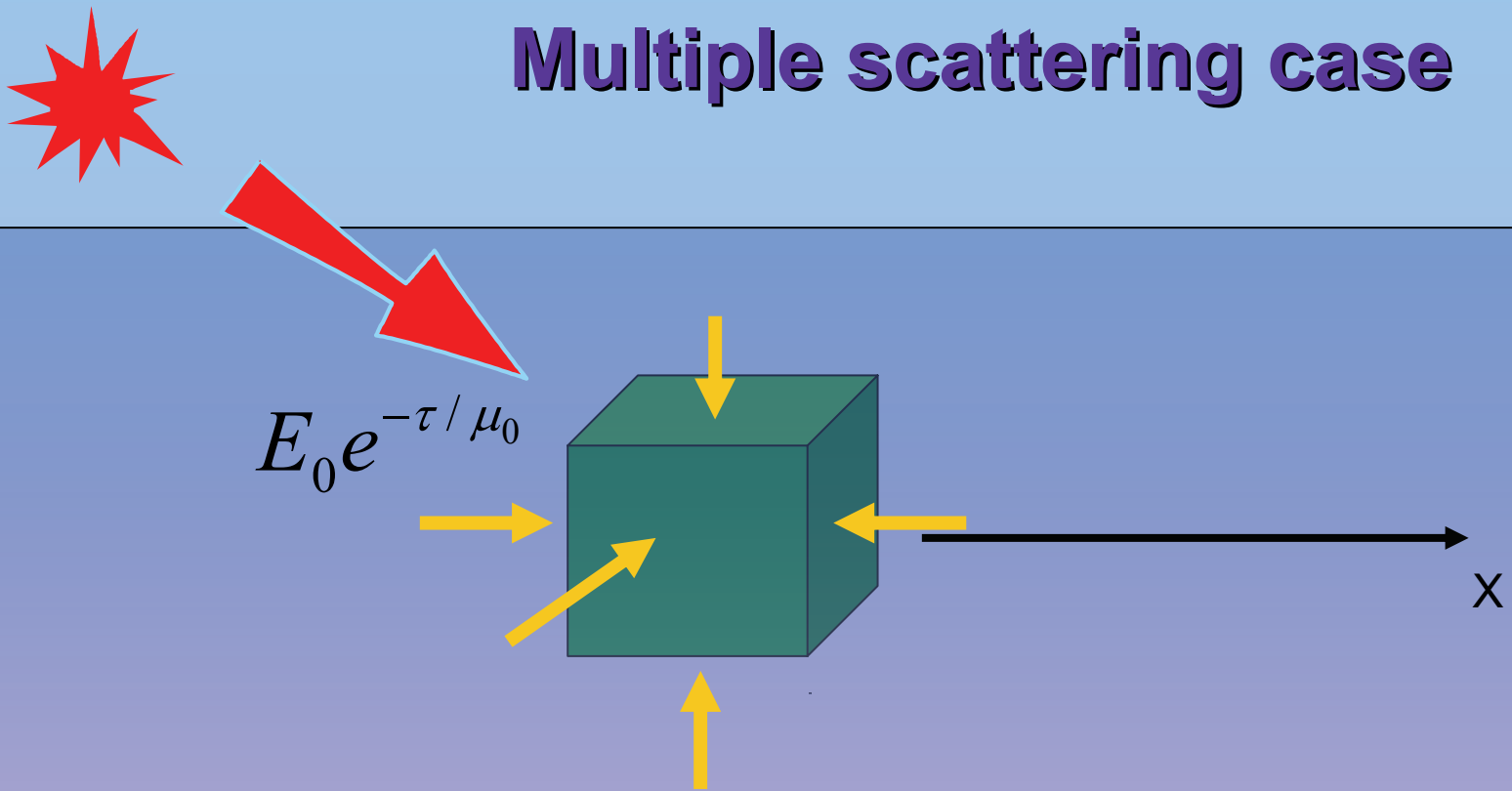


$$I_2 = \frac{E_0}{4\pi\mu} \cdot \omega_0 x(\mu) \int_0^\infty \alpha(z) \cdot e^{-\tau(z) \left(\frac{1}{\mu_0} + \frac{1}{\mu} \right)} dz$$

$$I_1 = \frac{E_0}{\pi} A_S \mu_0 e^{-\tau_0 \left(\frac{1}{\mu_0} + \frac{1}{\mu} \right)}$$

$$\mathcal{E} = E_0 e^{-\tau/\mu_0} \cdot \alpha \omega_0 \cdot \frac{x(\mu)}{4\pi}$$

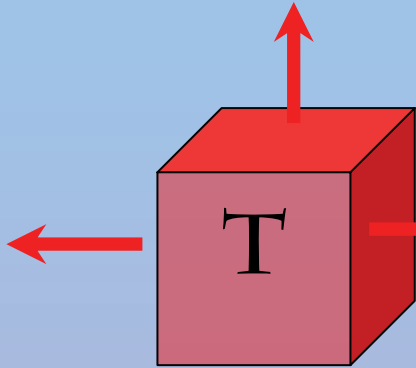
Multiple scattering case



$$\varepsilon = E_0 e^{-\tau/\mu_0} \omega_0 \alpha \frac{p(\gamma)}{4\pi} + \omega_0 \alpha \int_{\Omega} I(\omega') p(\gamma') \frac{d\omega'}{4\pi}$$

$$\frac{dI}{dx} = -\alpha I + \varepsilon$$

Thermal radiation



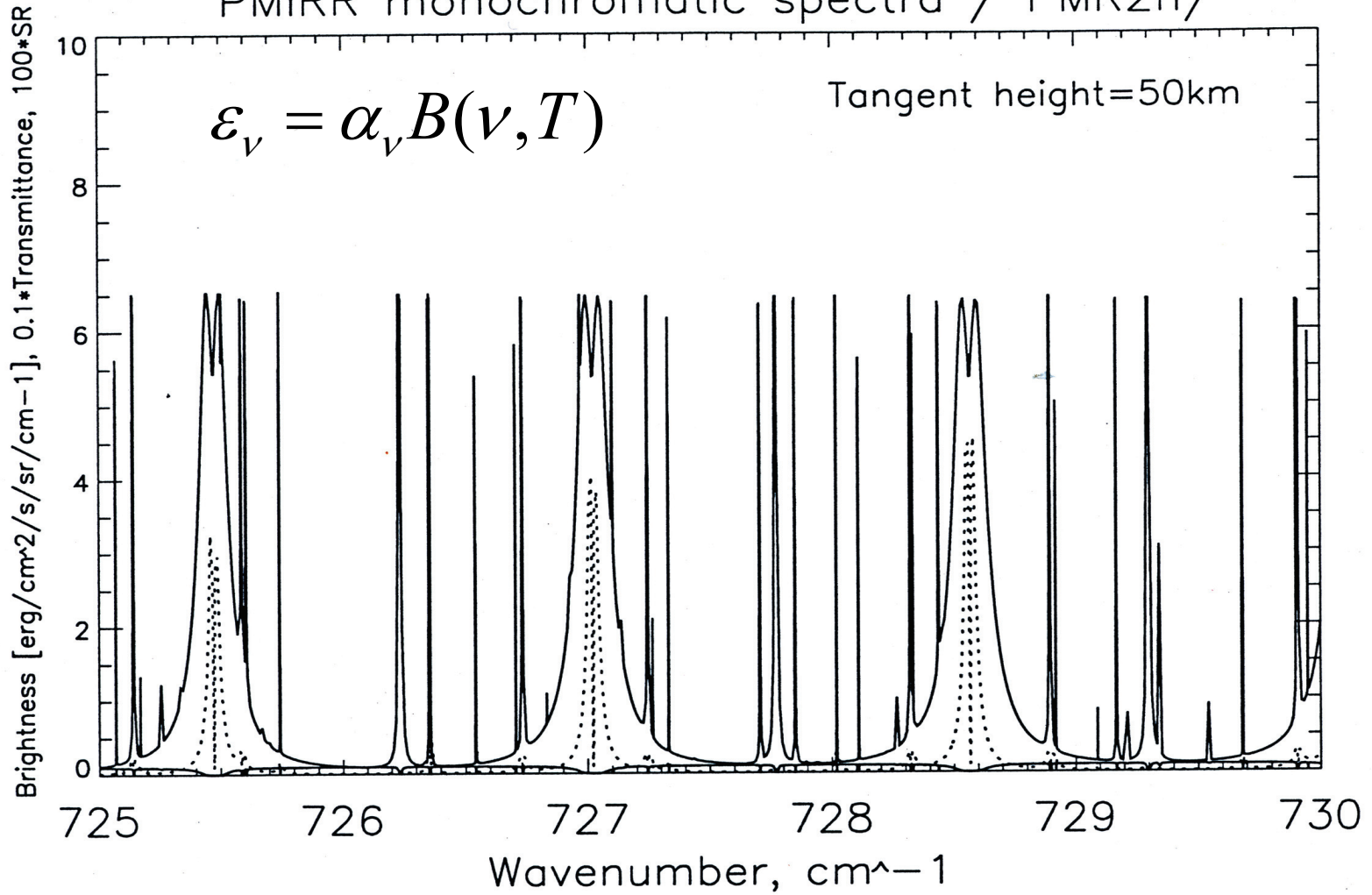
$$\varepsilon_\nu = \alpha_\nu B(\nu, T) = \alpha_\nu \frac{2h\nu^3}{c^2} \frac{1}{e^{h\nu/kT} - 1}$$

$$\Phi = \sigma T^4 \quad \text{- full thermal flux}$$

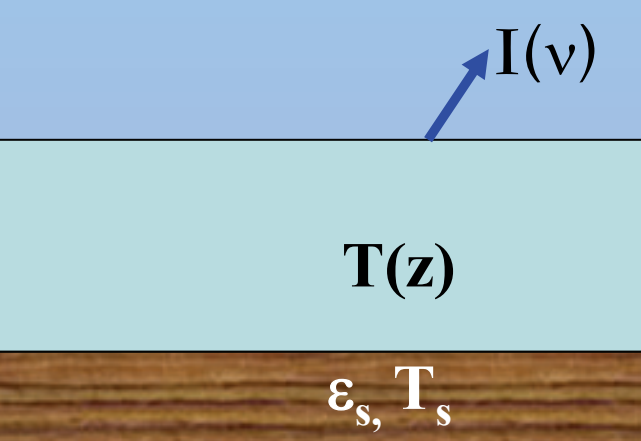
- ✚ Local thermodynamic equilibrium (LTE)
- ✚ LTE breakdown above relaxation level ($n \sim 10^{15} \text{ cm}^{-3}$)
- ✚ Thermal radiation in gaseous atmosphere

Atmospheric spectrum in CO₂ band

PMIRR monochromatic spectra / PMR2h/



Radiative transfer equation for thermal radiation



$$I(\nu) = \epsilon_s B_\nu(T_s)t + \int_{\text{Surface}}^{\text{Space}} B_\nu[T(z')] dt_\nu$$

Scattering is neglected

$$t_\nu(z') = \int_{z'}^{\text{space}} e^{-\tau_\nu(z)} dz - \text{transmittance}$$

$$I(\nu) = \epsilon_s B_\nu(T_s)t_\nu + \int_{\text{Surface}}^{\text{Space}} B_\nu[T(\xi)] \cdot K_\nu(\xi) d\xi$$

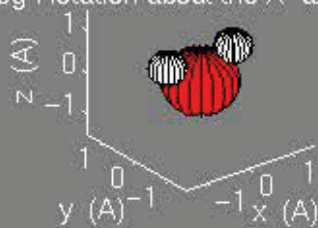
$$\xi = \lg p$$

$$K_\nu(\xi) = -\frac{\partial t_\nu}{\partial \xi} - \text{weighting function}$$

Basics of spectroscopy

Molecular "gymnastics"

Bond length = 0.96 Å Bond angle = 105 deg Rotation about the X-axis



Rotation about the Y-axis



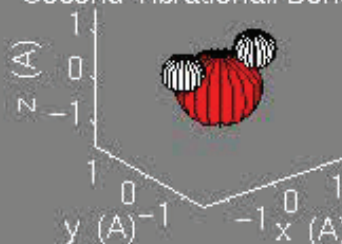
Rotation about the Z-axis



Bond length=0.96Å angle=105deg First Vibrational: Symmetric Stretching



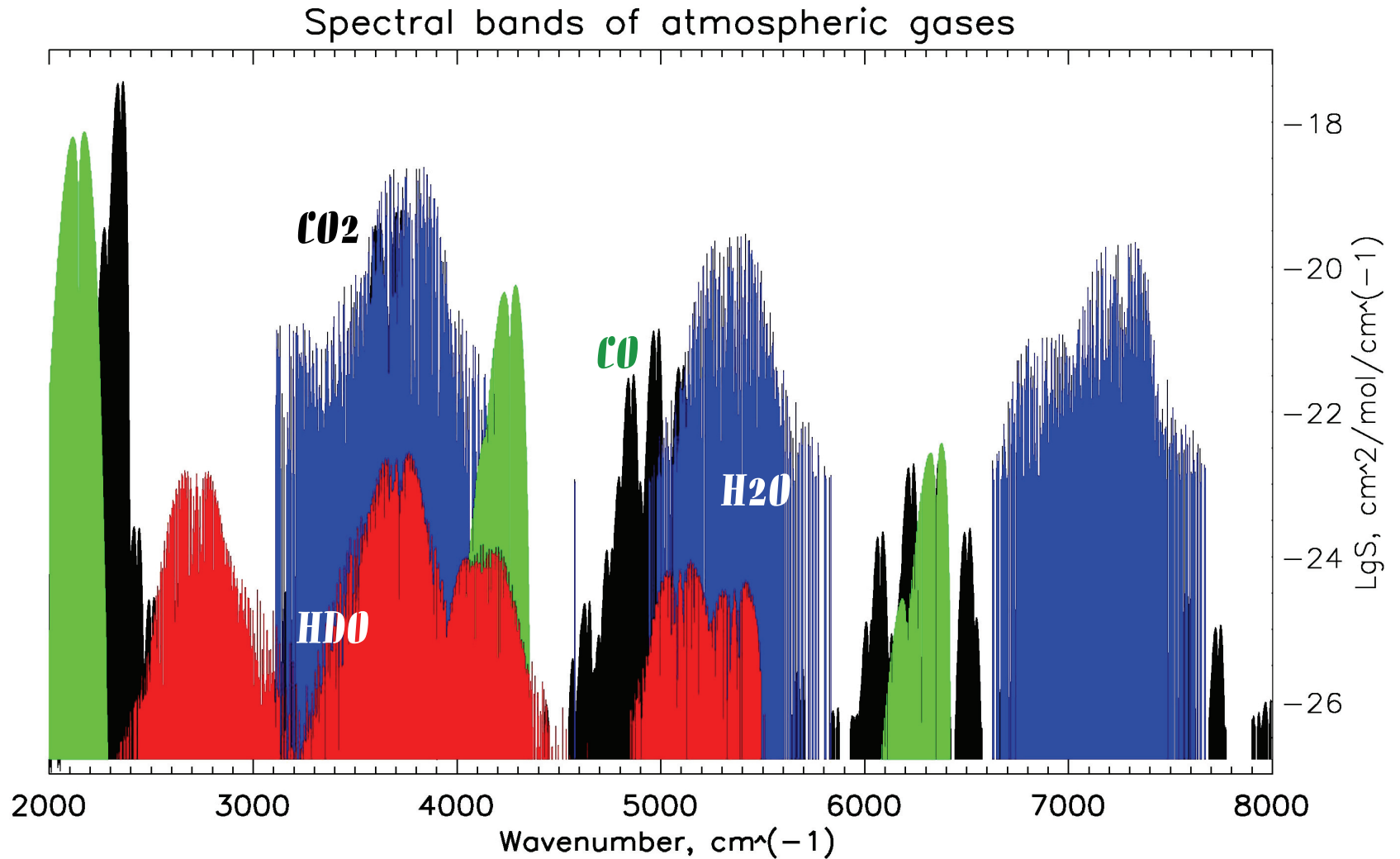
Second Vibrational: Bending



Third Vibrational: Asymmetric Stretching



Spectral bands of atmospheric gases



Spectral line shapes

1. Spectral line shape: $k(\nu) = S \cdot L(\nu - \nu_0)$

2. Natural broadening: $\Delta\nu \sim 10^{-10} \text{ cm}^{-1}$

3. Doppler broadening

$$L_D = \frac{1}{\Delta\nu \sqrt{\pi}} \cdot e^{-\left(\frac{\nu - \nu_0}{\Delta\nu}\right)^2}$$

$$\Delta\nu = \frac{\nu_0}{c} \sqrt{\frac{3kT}{m}} \approx 10^{-2} - 10^{-4} \text{ cm}^{-1}$$

4. Lorentz broadening

$$L_L = \frac{1}{\pi} \frac{\alpha_L}{(\nu - \nu_0)^2 + \alpha_L^2}$$

$$\alpha_L = \alpha_0 \frac{P}{P_0} \left(\frac{T_0}{T}\right)^2$$

5. Voigt line profile

$$L_V = \frac{\sqrt{\ln 2} a}{\pi^{3/2} \Delta\nu} \int_{-\infty}^{\infty} \frac{e^{-y^2} dy}{(x-y)^2 + a^2}$$

$$x = \frac{(\nu - \nu_0) \sqrt{\ln 2}}{\Delta\nu}$$

$$a = \frac{\alpha_L \sqrt{\ln 2}}{\Delta\nu}$$

Solar range: Equivalent width and curve of growth

1. Equivalent width
2. Absorption in isolated Lorentz line

$$W(N) = 2\pi d_L \cdot L(u) \begin{matrix} \rightarrow \sim u, u \ll 1 \\ \rightarrow \sim \sqrt{u}; u \gg 1 \end{matrix}$$

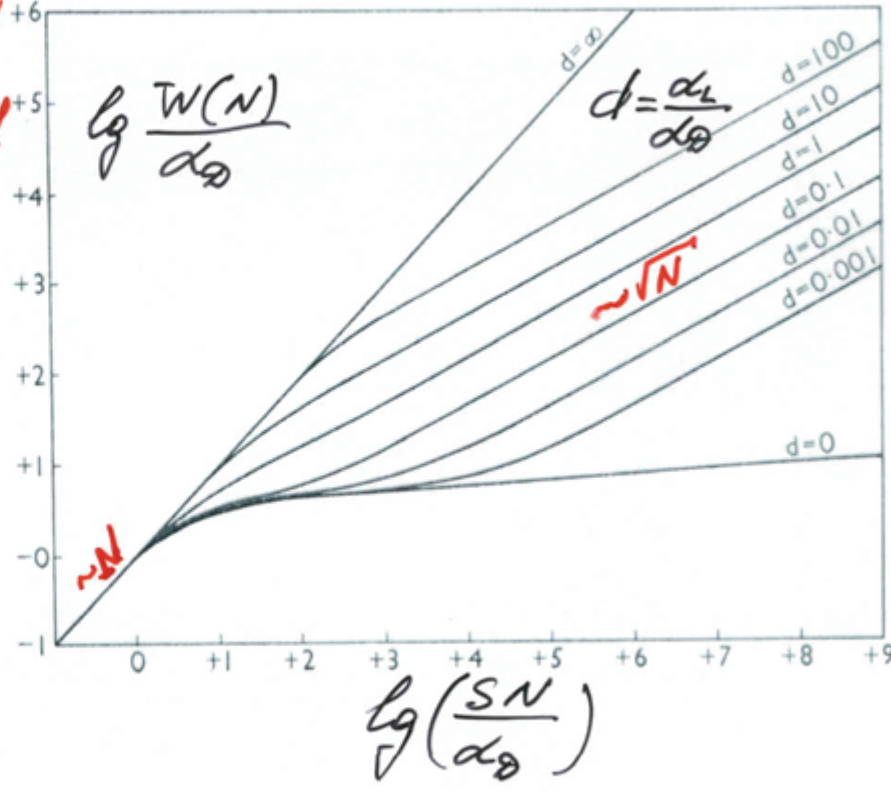
$$u = \frac{SN}{2\pi d_L}$$

3. Absorption in isolated Voght line

4. Elsasser band model

$$Z(x) = 2\pi u y \frac{\sin 2\pi y}{\text{ch } 2\pi y - \cos 2\pi x}$$

$$u = \frac{SN}{2\pi d_L}, \quad x = \frac{\nu}{\delta}, \quad y = \frac{d_L}{\delta}$$



Solar reflected spectrum depends mainly on gas amount and to a less extent on the temperature!

Spectra of thermal emission

1. Radiative transfer equation

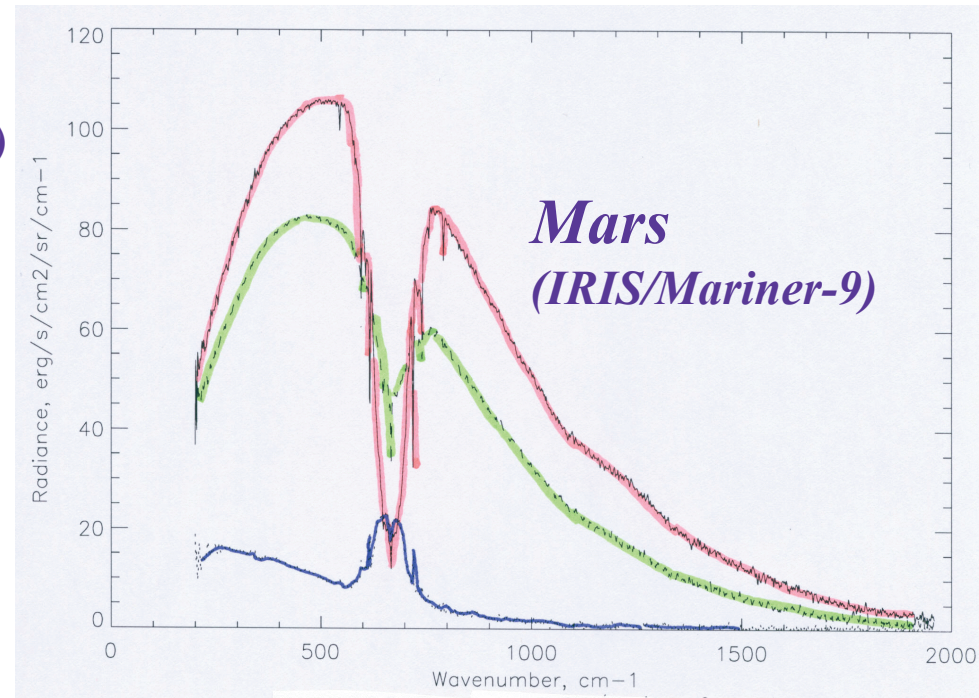
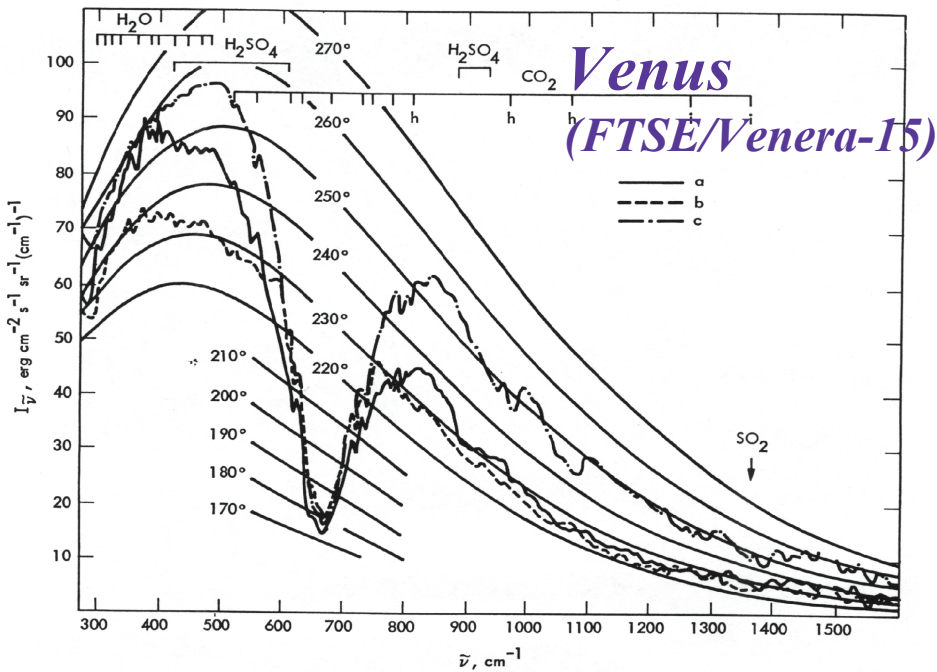
$$I = \epsilon_s \cdot B(T_s) \cdot t + \int_0^t B[T'(s')] ds'$$

$$I = \epsilon_s B(T_s) t + \int_{-\infty}^{\infty} B[T(\xi)] K(\xi) d\xi \quad \xi = \log p$$

2. Weighting function

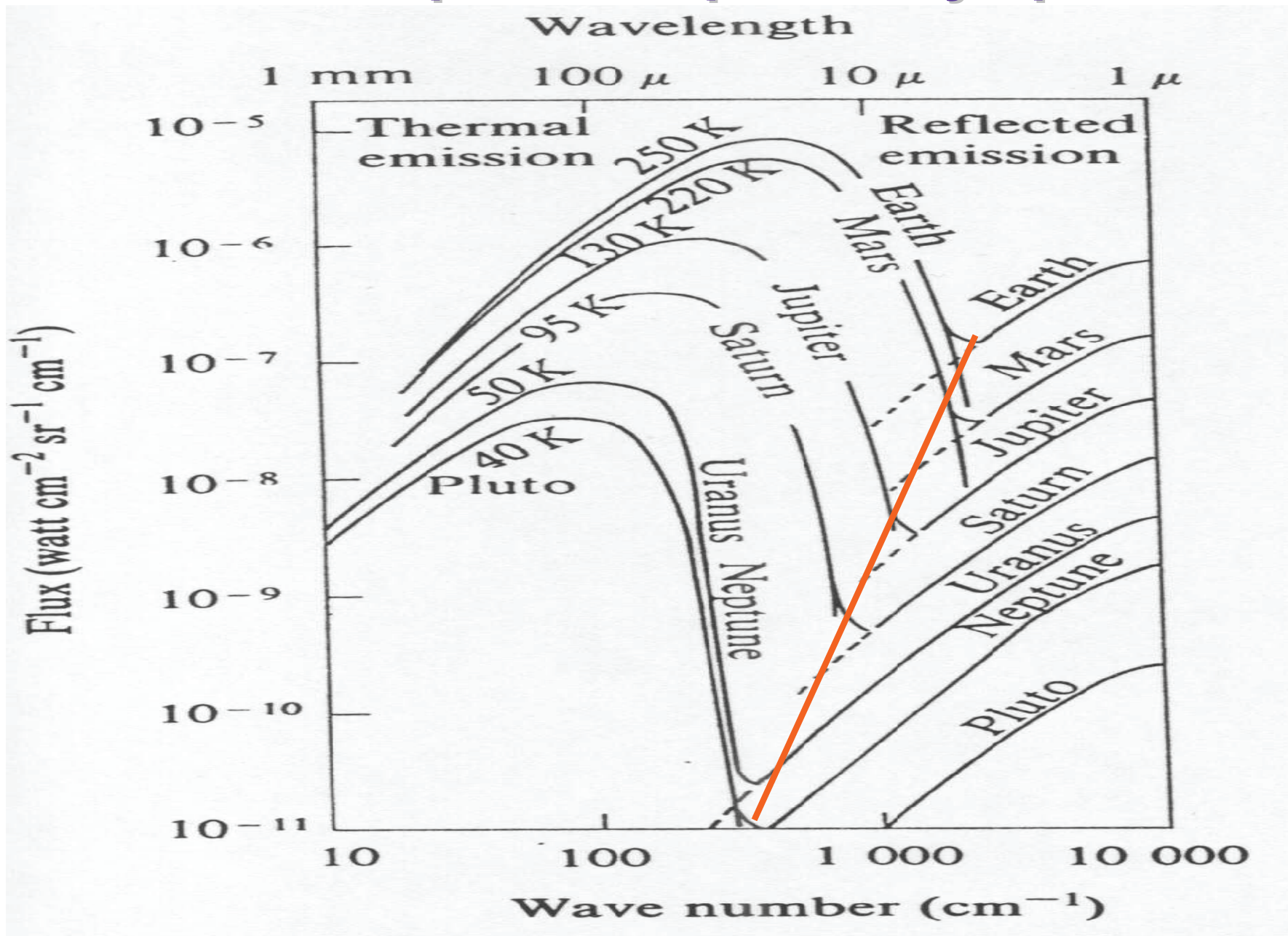
$$K(\xi) = -\frac{\partial t}{\partial \xi}$$

3. Examples of thermal spectra of planets



Thermal spectrum depends temperature and gas abundance!

General shape of the planetary spectra





*Additional slides to
Basics of Radiative Transfer*

Principles of temperature sounding

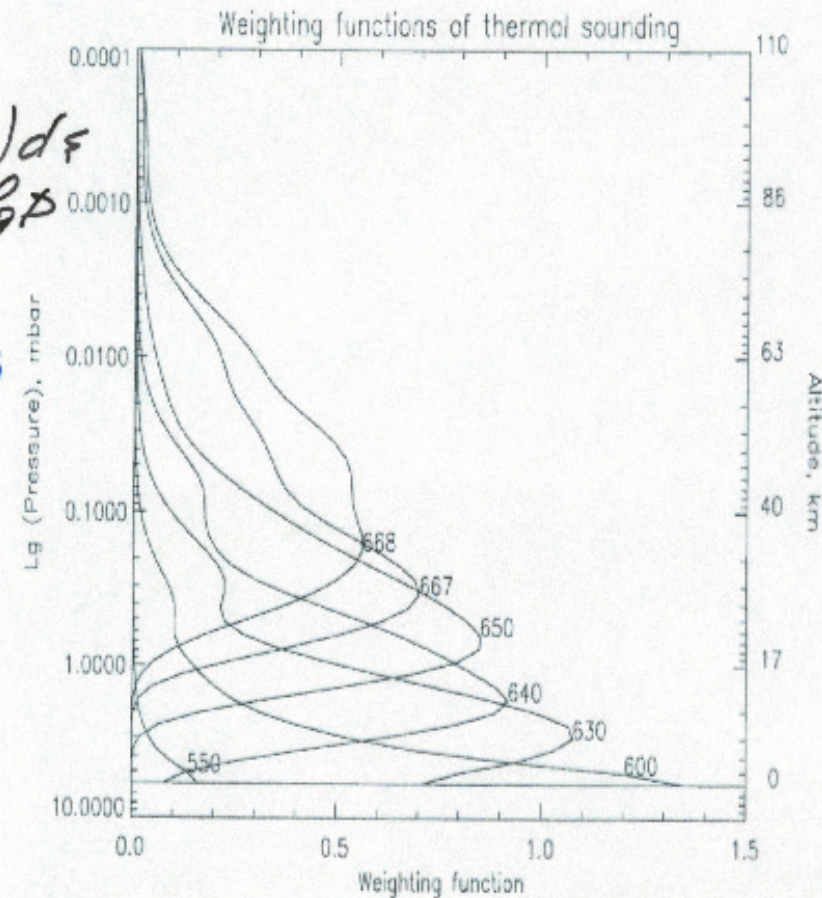
- In strong bands of atmospheric gases thermal radiation forms at different altitudes depending on wavelength

- $$I(\nu) = \epsilon_s B(T_s) \tau_{ATM} + \int_{-\infty}^{\infty} B[T(\xi)] \cdot K(\xi) d\xi$$

$$K(\xi) = -\frac{dT}{d\xi} \quad \xi = \ln p$$

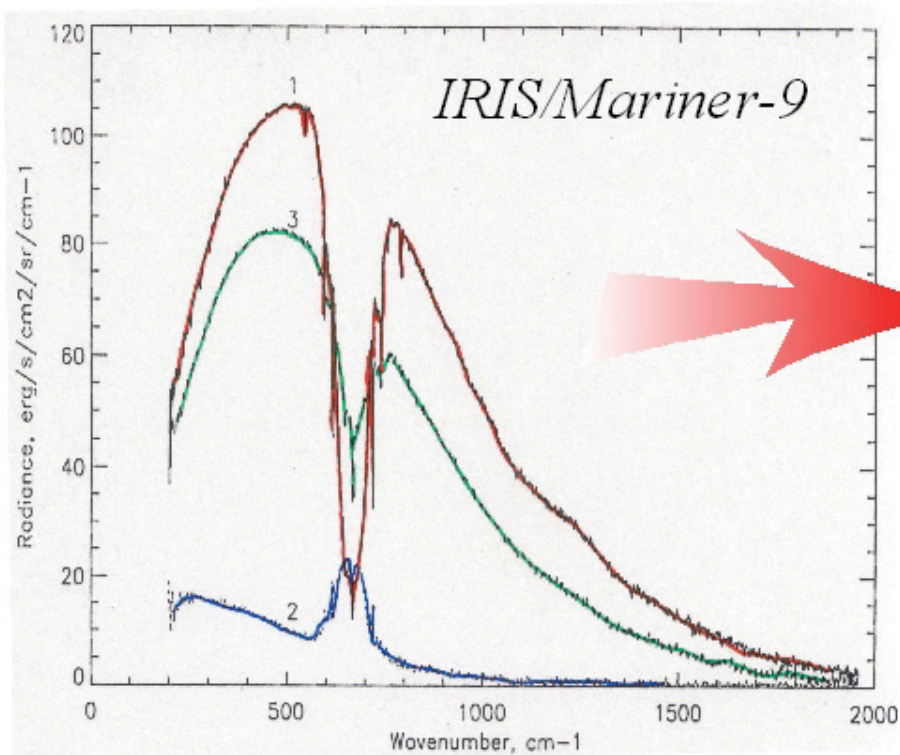
- Gas should be well mixed and its abundance known
- Local thermodynamic equilibrium
- Temperature retrieval is an **ill-posed** problem

Weighting functions of thermal sounding at Mars

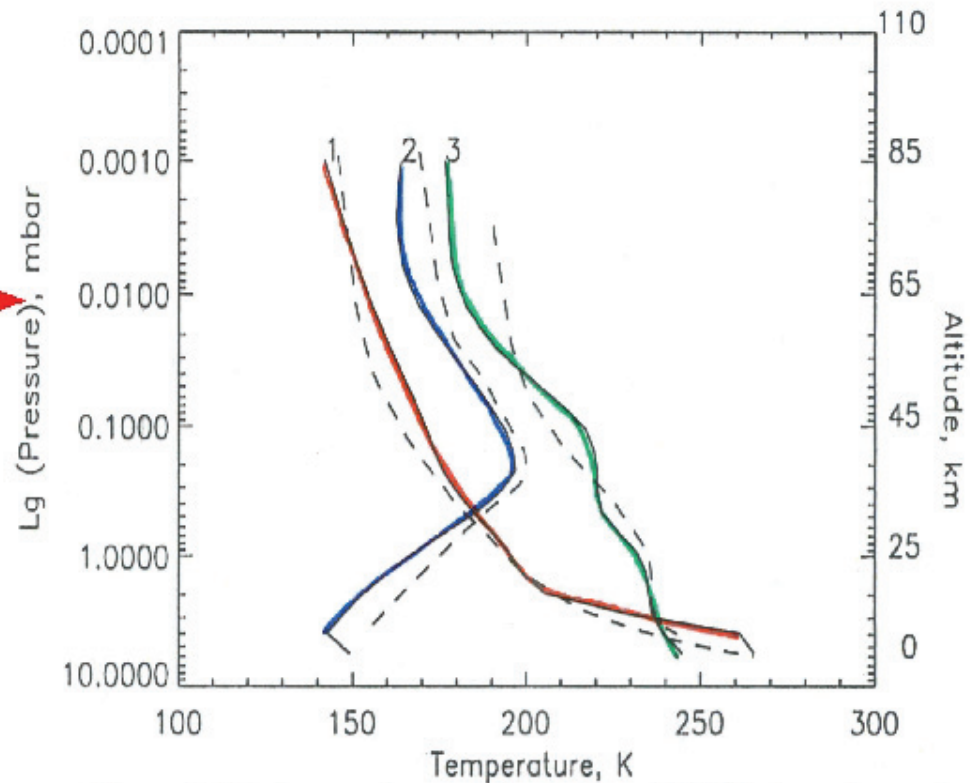


Example of temperature sounding

Measured spectra of Mars



Retrieved temperature profiles



Scattered solar light in the atmosphere (2)

5. Two stream approximation /Schwarzschild- Schuster/

$$I^\uparrow(\tau) = \int I(\tau) \sin\theta d\theta \quad I^\downarrow(\tau) = \int I(\tau) \sin\theta d\theta$$

$$\begin{cases} \frac{1}{2} \frac{dI^\uparrow}{d\tau} = I^\uparrow - S \\ -\frac{1}{2} \frac{dI^\downarrow}{d\tau} = I^\downarrow - S \end{cases} \Rightarrow S(\tau) = F\left(\tau + \frac{1}{2}\right)$$
$$S(\tau) = \frac{1}{2} (I^\uparrow + I^\downarrow)$$

6. N-stream approximation

$$\mu_i \frac{dI(\tau, \mu_i)}{d\tau} = I(\tau, \mu_i) - \frac{1}{2} \sum_j a_j I(\tau, \mu_j) \quad i = \pm 1 \dots \pm N$$

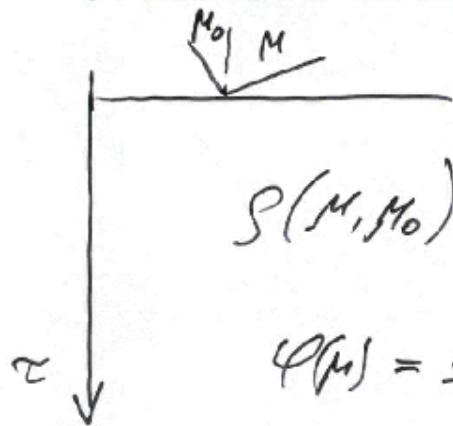
Scattered solar light in the atmosphere (3)

3. Spherical phase function $X(\gamma)=1$ /Hvolson/

$$\begin{cases} \mu \frac{dI}{d\tau} = I - S \\ S = \frac{\omega_0}{2} \int_{-1}^1 I(\mu) d\mu + \frac{\omega_0}{4\pi} E_0 e^{-\tau/\mu_0} \end{cases}$$

$$\Rightarrow S(\tau) = \frac{\omega_0}{2} \int_0^{\infty} E_{i_1}(\tau-t) S(t) dt + \frac{\omega_0}{4\pi} E_0 e^{-\tau/\mu_0}; \quad E_{i_1}(x) = \int_1^{\infty} \frac{e^{-xy}}{y} dy$$

4. Reflection from semi-infinite atmosphere /Ambartsumjan/

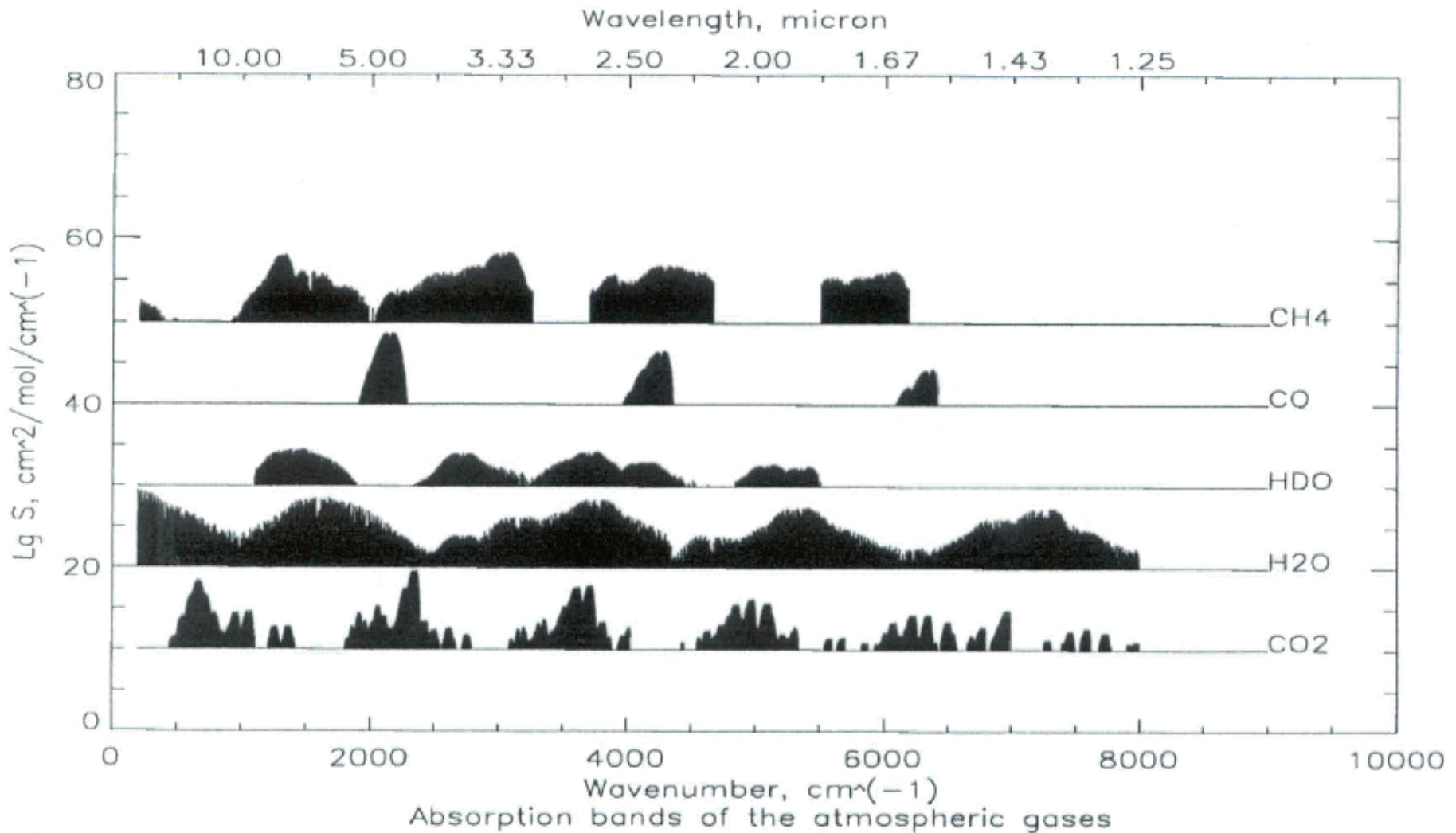


$$S(\mu, \mu_0) = \frac{\omega_0}{4} \frac{\varphi(\mu) \cdot \varphi(\mu_0)}{\mu + \mu_0}$$

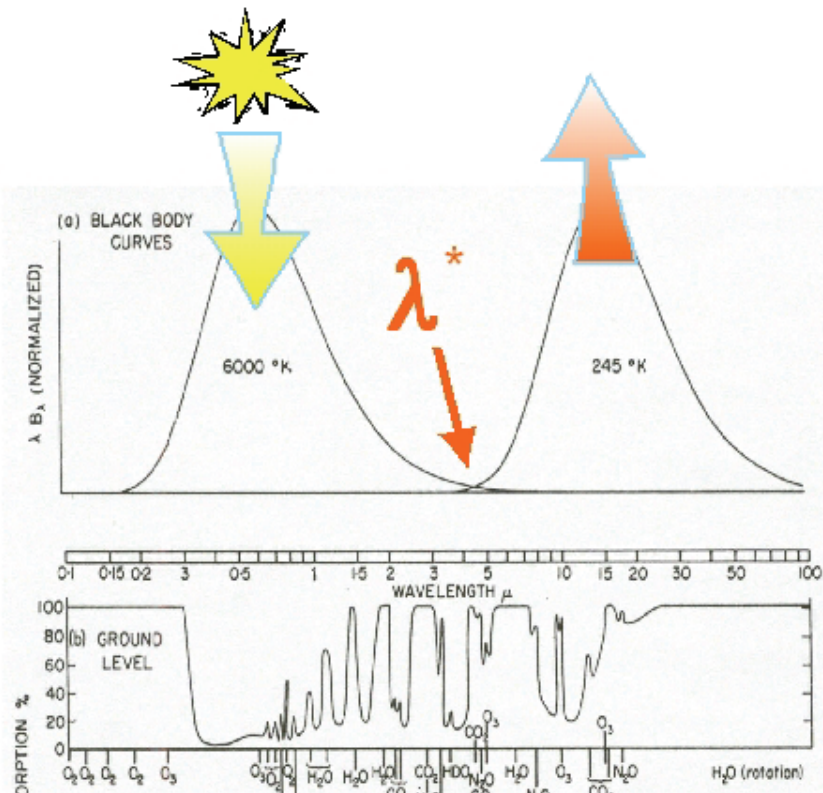
$$\varphi(\mu) = 1 + \frac{\omega_0}{2} \varphi(\mu) \int_0^1 \frac{\varphi(\mu')}{\mu + \mu'} d\mu'$$

Basics of spectroscopy (3)

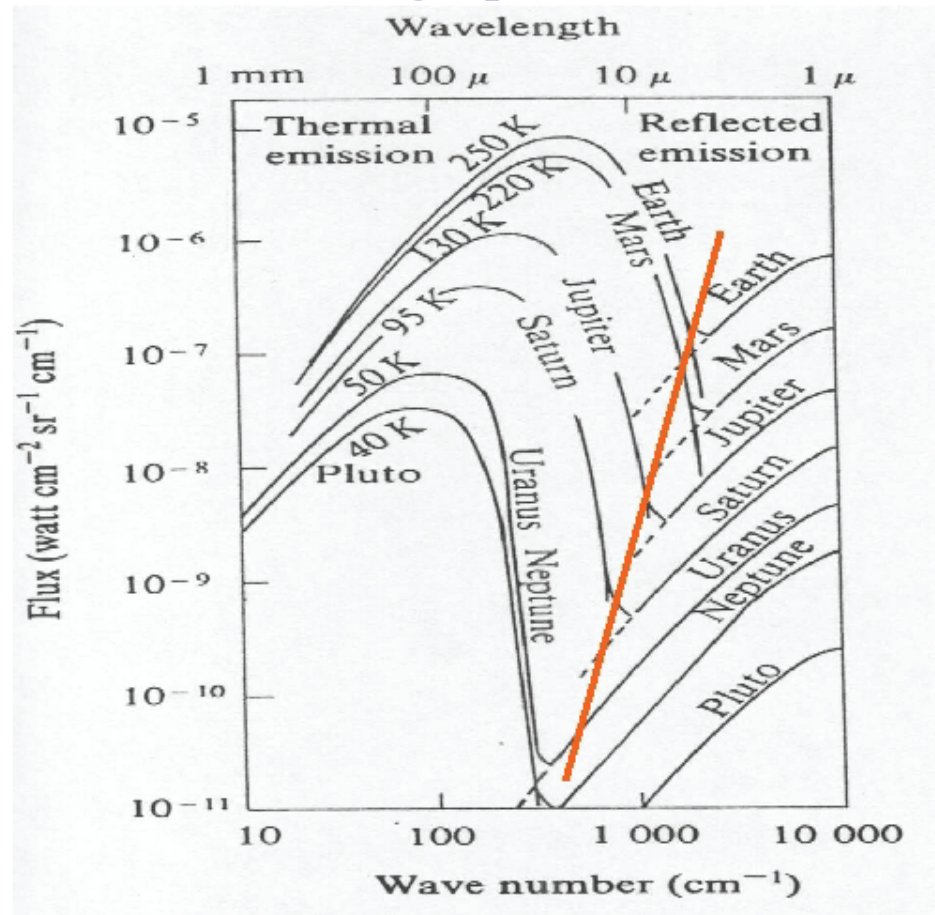
Absorption bands of atmospheric gases



Energy balance of planets



Planetary spectra



Effective temperature

$$E(1-A) = 4\sigma T_{\text{eff}}^4$$

Solar constant

$$E_{\oplus} = 1370 \text{ W/m}^2$$